

Ocean waves:

We are dealing with wind generated waves from gentle to rough ...

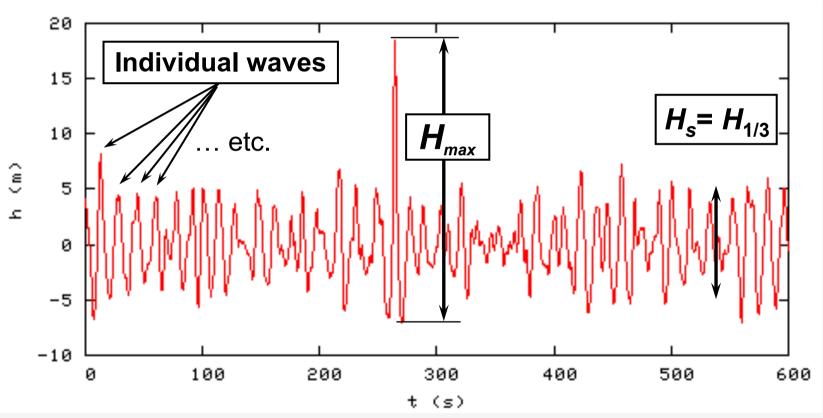


February 5, 2014

Observe Individual Waves,

After a while, you can estimate a characteristic height the waves: the Significant Wave Height, H_s ,

You might also notice that some waves are larger than the rest, characterised by the Maximum Individual Wave Height, H_{max}

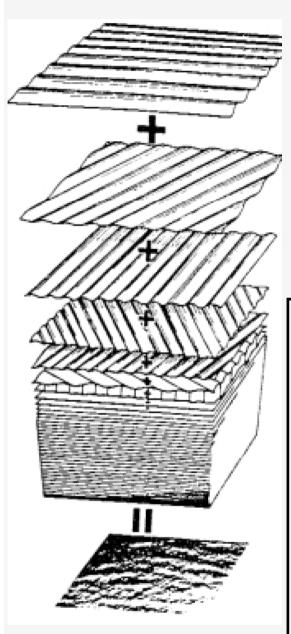


Surface elevation time series from platform Draupner in the North Sea

How do we go about making predictions on the sea state?



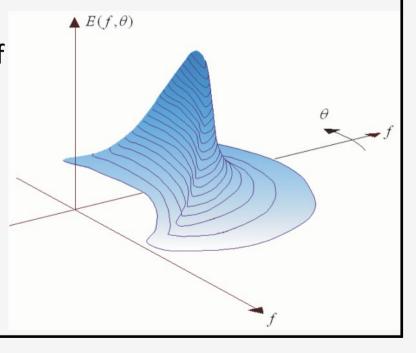




Wave Spectrum

• The irregular water surface can be decomposed into (*infinite*) number of simple sinusoidal components with different frequencies (f) and propagation directions (θ).

The distribution of wave energy among those components is called:
 "wave spectrum", F(f, θ).

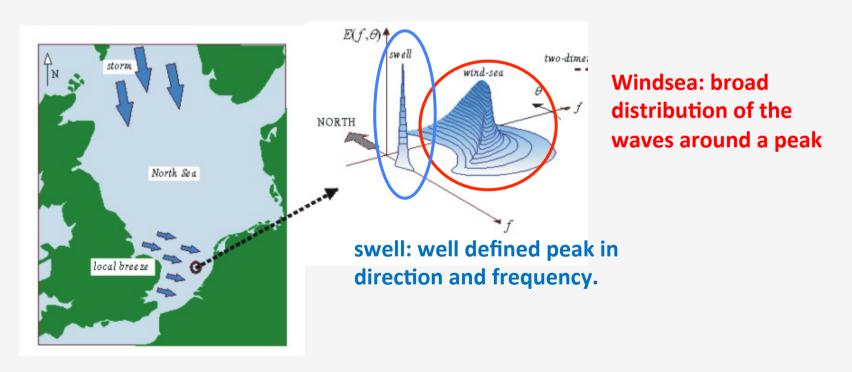


Modern ocean wave prediction systems are based on statistical description of oceans waves (i.e. ensemble average of individual waves).

The sea state is described by the two-dimensional wave spectrum $F(f, \theta)$.

For instance, the sea state off the coast of Holland might the results of a local sea breeze. These waves are generally known as windsea

Waves might have also propagated from their generation area as swell



Ocean Wave Modelling

The 2-D spectrum follows from the energy balance equation (in its simplest form: deep water case, no surface currents):

$$\frac{\partial F}{\partial t} (\vec{V_g}) \nabla F = S_{in} + S_{nl} + S_{diss}$$

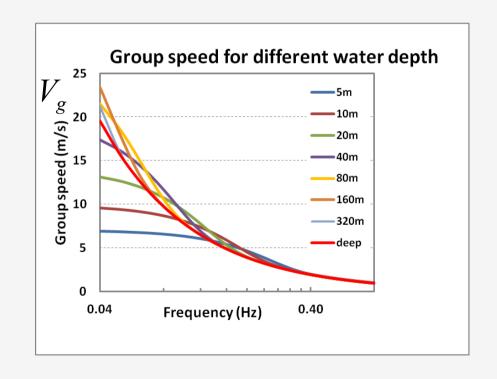
Where the group velocity V_g is derived from the dispersion relationship which relates frequency (f) and wave number (k) for a given water depth (D).

$$\omega^2 = g k \tanh(kD)$$

$$V_g = \frac{\partial \omega}{\partial k}$$
$$\omega = 2\pi f$$

$$\omega = 2\pi f$$

D: water depth



Ocean Wave Modelling

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Where the group velocity V_g is derived from the dispersion relationship which

relates frequency (f) and wave number (k) for a given water depth (D).

S_{in}: wind input source term (generation).

S_{nl}: non-linear 4-wave interaction (redistribution).

tings wind-induced air pressure, the wind-induced air pressure, surface moting discovered.

Figure 6.16 The wave-induced wind-pressure variation over a propagating harmonic wave.

The wave grows by this mechanism, the mechanism becomes the wave can therefore grow faster, which in turn makes the ms effective, etc.

Figure 6.00 Density-pie wave-wave interactions (real-fields in indeep votate). Two pieces of wave components can create most dismond patterns with desired wave lengths and directions and therefore discribed when the four waves are superimposed (not shown here), they can thus sessonate. The wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand panel in wave-number vectors of the four wave components are shown in the right-hand pan

S_{diss}: dissipation term due to whitecapping (dissipation).





• Once you know the wave spectrum F, any other sea state parameters can be estimated. For example, the mean variance of the sea surface elevation η due to waves is given by:

$$\langle \eta^2 \rangle = \iint F(f, \theta) df d\theta$$

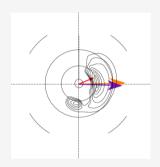
• The statistical measure for wave height, called the significant wave height (H_s) :

$$H_{S} = 4\sqrt{\langle \eta^{2} \rangle}$$

The term significant wave height is historical as this value appeared to be well correlated with visual estimates of wave height from experienced observers.

It can be shown to correspond to the average $1/3^{rd}$ highest waves $(H_{1/3})$.

2-D spectrum can be used to specify boundary conditions for limited area wave model.

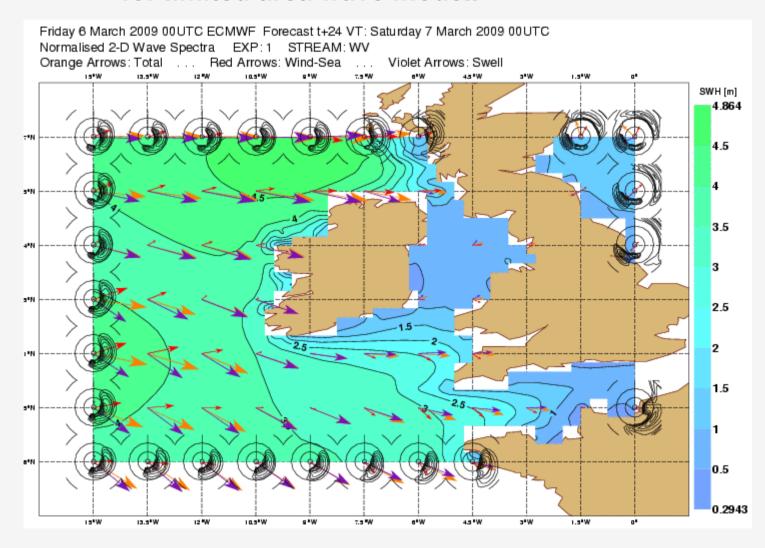


Spectra used as boundary conditions

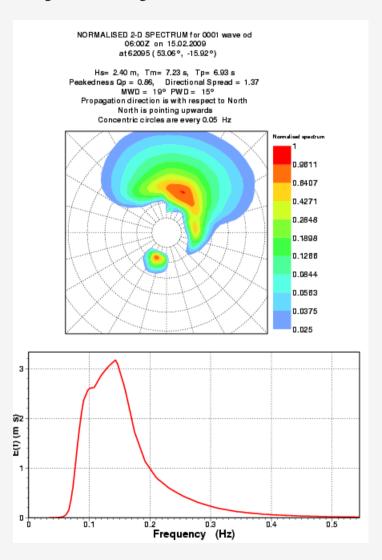
→ windsea

→ swell

→ total sea



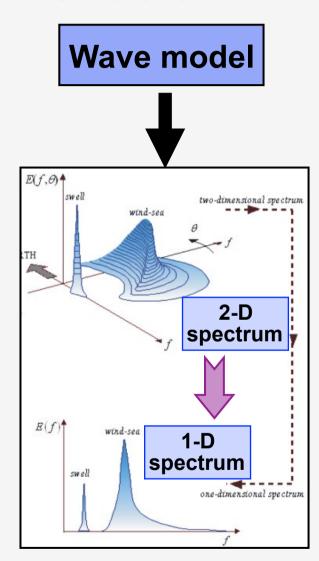
Or if you only look at one location ...



The complete description of the sea state is given by the 2-D spectrum, however, it is a fairly large amount of data.

It is therefore reduced to integrated quantities:

1-D spectrum obtained by integrating the 2-D spectrum over all directions and/or over a frequency range.



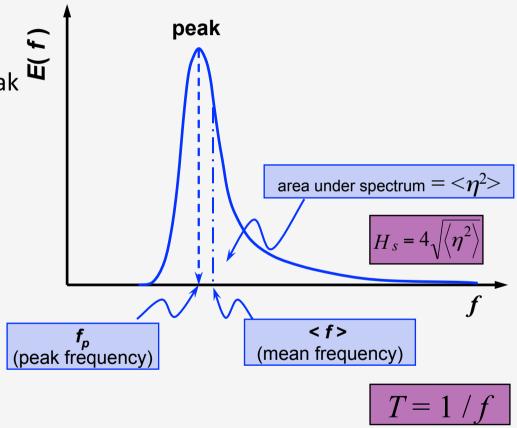
When simple numbers are required, the following parameters are available:

The significant wave height (H_s).

The peak period (period of the peak of the 1-D spectrum).

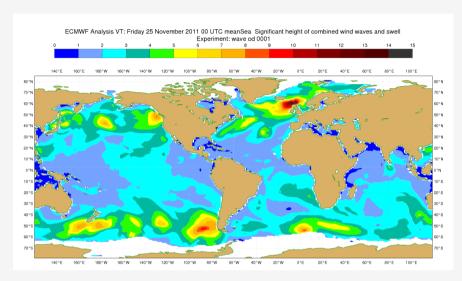
Mean period(s) obtained from weighted integration of the 2-D spectrum.

Integrated mean direction. and many others.



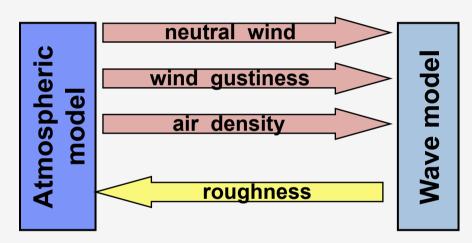
Complete list at: http://www.ecmwf.int/services/archive/d/parameters/order=/table=140/

ECMWF Wave Model Configurations



Global from 81°S to 90°N

Coupled to the atmospheric model with feedback of the sea **surface roughness** change due to waves.



ECMWF Wave Model Configurations

High resolution (HRES-WAM)

- 14 km grid spacing.
- Dissemination grid: 0.125°x0.125°
- 36 frequencies.
- 36 directions.
- Coupled to the TCo1279 model.
- Analysis every 6 hrs and 10 day forecasts from 0 and 12 UTC.

Ensemble forecasts (EPS-WAM)

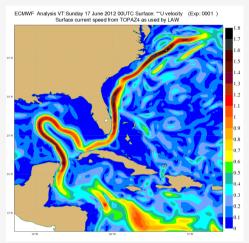
- 28 km grid spacing.
- Dissemination grid: 0.25°x0.25°
- 36 frequencies.
- 36 directions.
- Coupled to TCo639.
- (50+1) 15 day forecasts from 0 and 12UTC (monthly twice a week).
- Coupled to ocean model.

ECMWF Wave Model Configurations

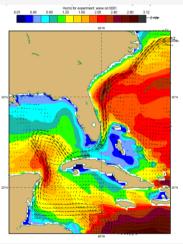
We also have the uncoupled global configuration (stand alone)

High resolution stand alone(HRES-SAW)

- 11 km grid spacing.
- Dissemination grid: 0.1°x0.1°
- 36 frequencies.
- 36 directions.
- Forced by HRES fields.
- Analysis every 6 hrs and 10 day forecasts from 0 and 12 UTC.
- Imposed surface currents from TOPAZ4 system.



TOPAZ4 surface currents

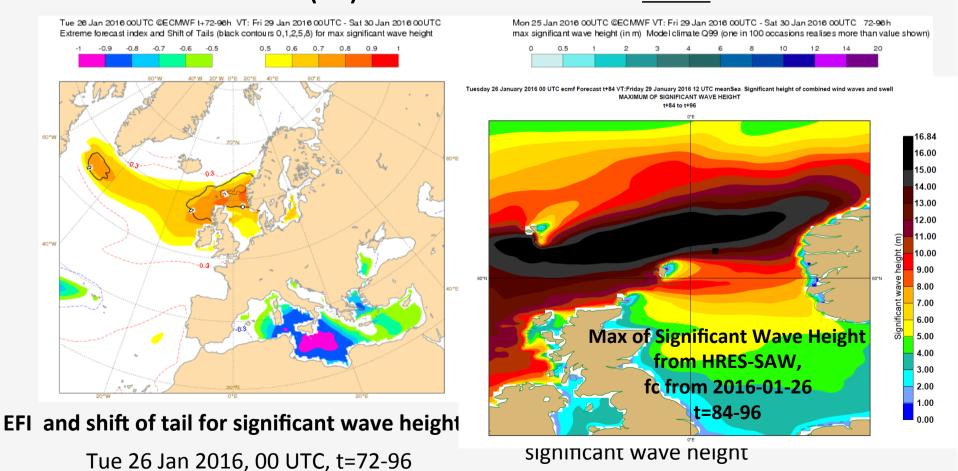


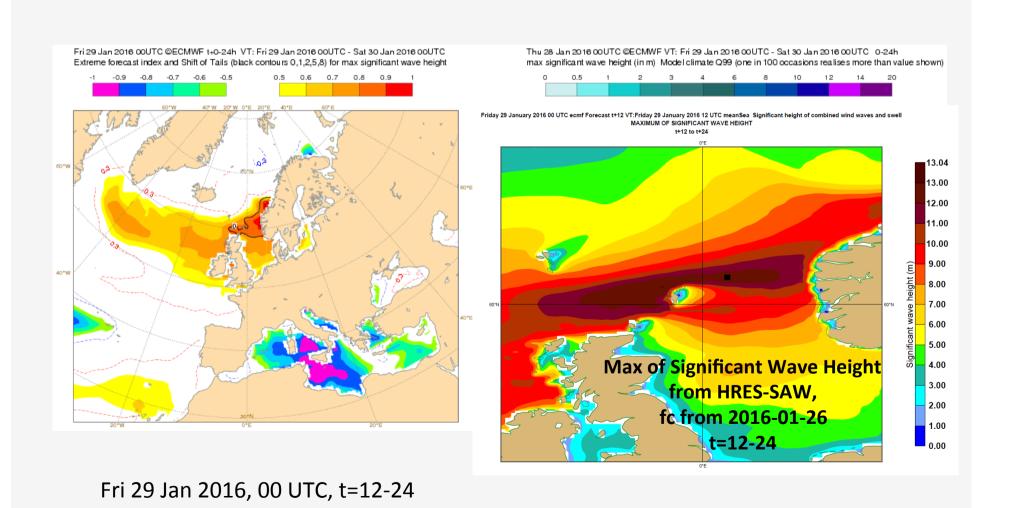
Wave Model Products: EFI plots

From the model climate, it is possible to derive indices that indicate deviations in probabilistic terms from what is 'expected'.

Extreme Forecast Index (EFI): 1 means that all EPS are <u>above</u> climate.

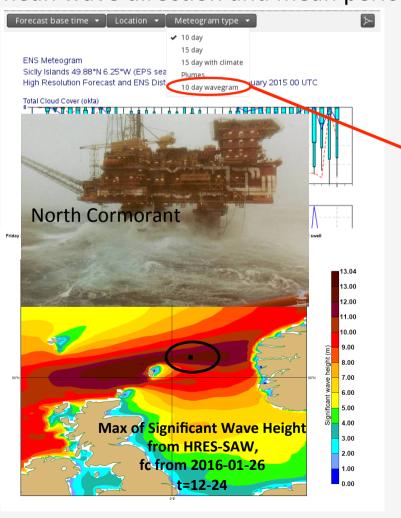
Gertrude (UK), Tor (Norway)

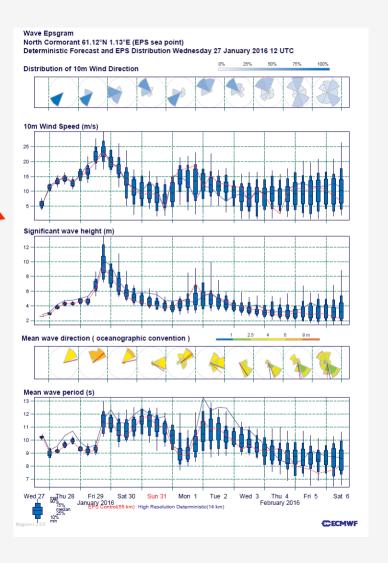




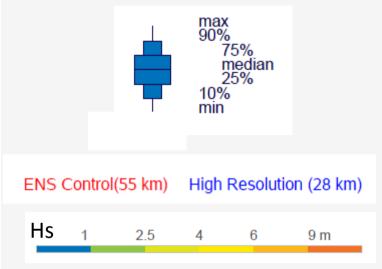
A bit more compact: Wave EPSgram:

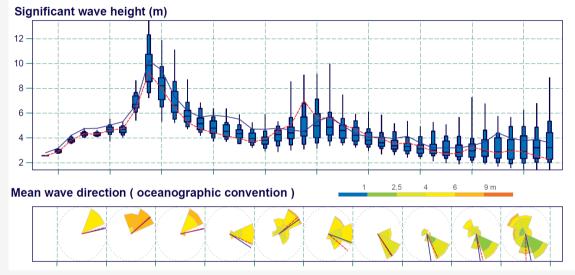
Like normal EPSgram but for wind direction, wind speed, significant wave height, mean wave direction and mean period.

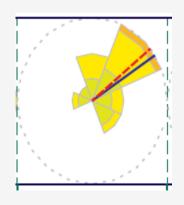




A bit more compact: Wave EPSgram:



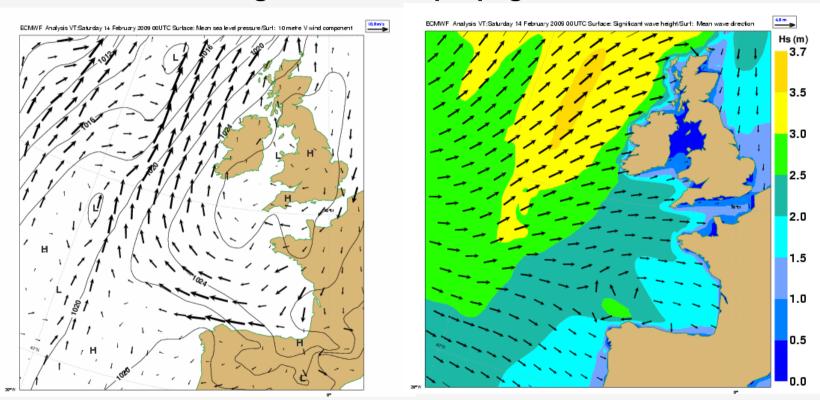




Each octant is coloured based on the distribution of the significant wave height associated with each mean direction. The coloured areas correspond to the fractional number of ensemble members with wave height in the range specified by the coloured ruler.

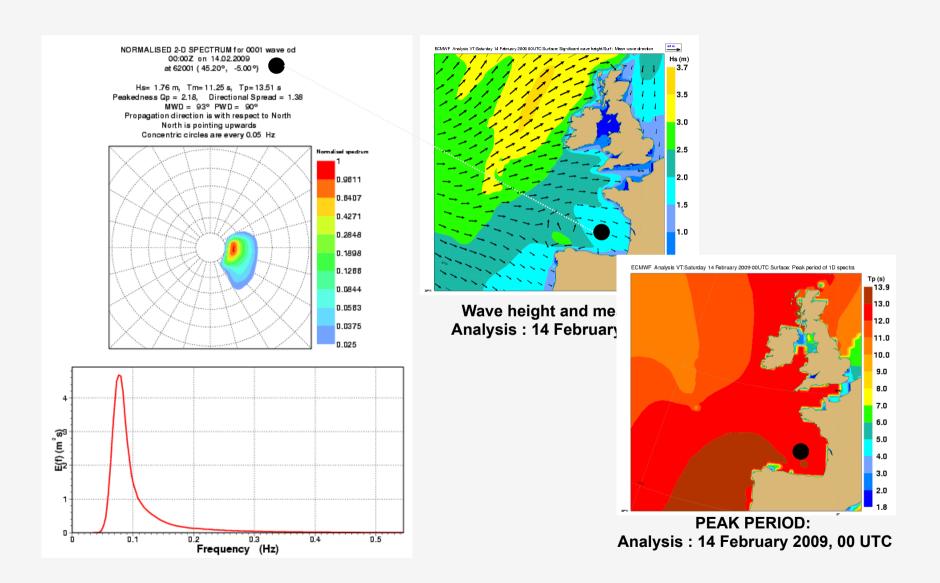
A bit more on Wave Model Products

Use simple parameters: total wave height and mean propagation direction

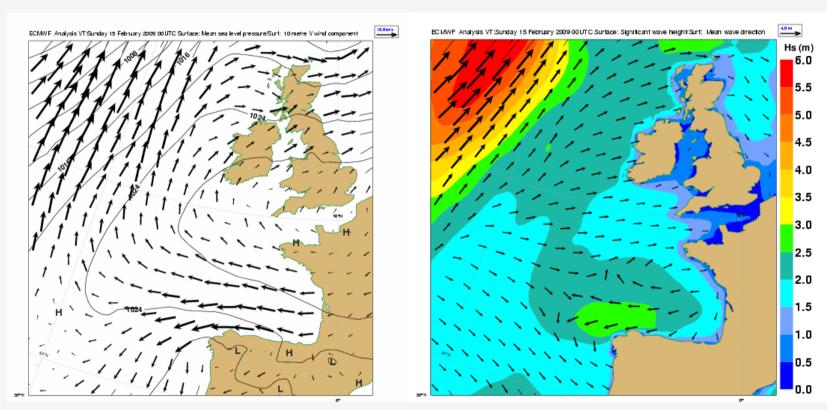


10m winds and mean sea level pressure: Analysis: 14 February 2009, 00 UTC

Wave height and mean direction: Analysis: 14 February 2009, 00 UTC

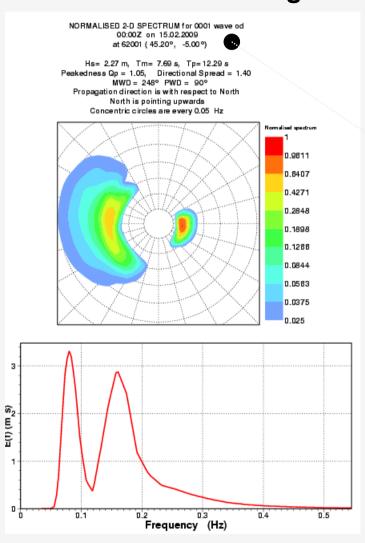


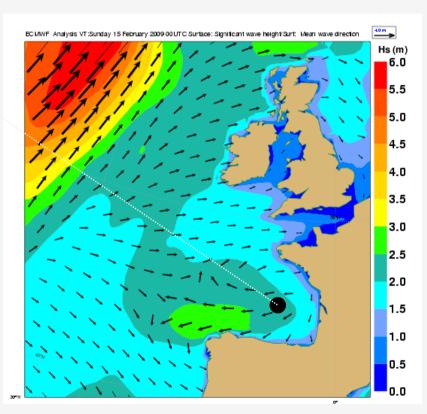
Situation might be more complicated!



10m winds and mean sea level pressure: Analysis: 15 February 2009, 00 UTC Wave height and mean direction: Analysis: 15 February 2009, 00 UTC

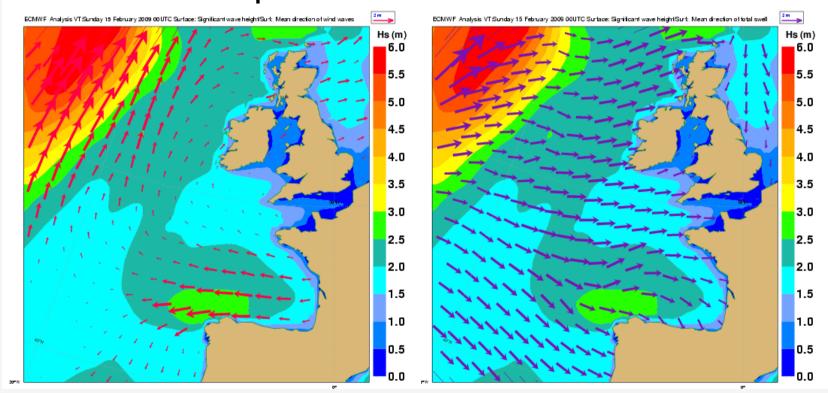
Situation might be more complicated:





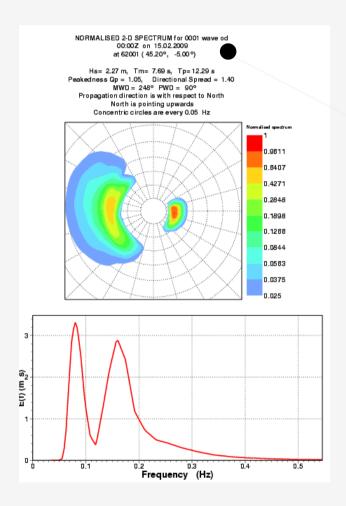
Wave height and mean direction: Analysis: 15 February 2009, 00 UTC

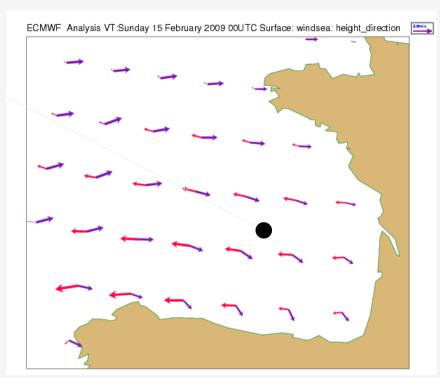
A scheme is used to split the global wave fields into waves which are under the direct influence of the forcing wind, the so-called windsea or wind waves, and those waves that are no longer bound to the forcing wind, generally referred to as swell. Period and mean direction are also determined for these split fields.



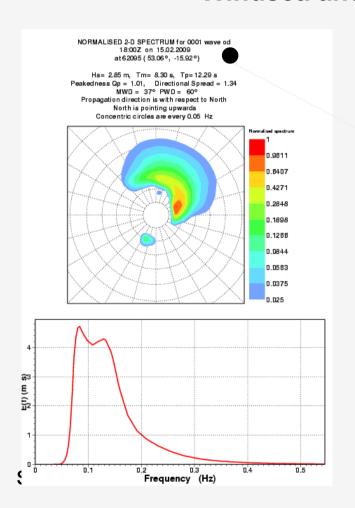
Wave height and windsea mean direction: Analysis: 15 February 2009, 00 UTC Wave height and swell mean direction: Analysis: 15 February 2009, 00 UTC

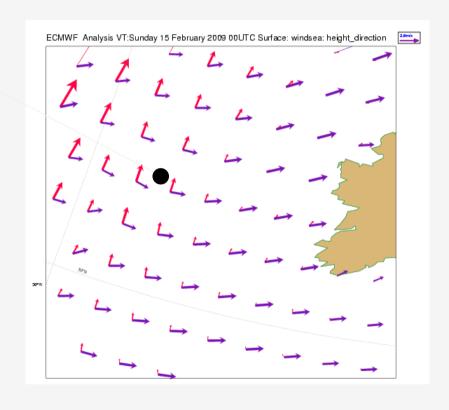
Windsea and swell: opposing sea



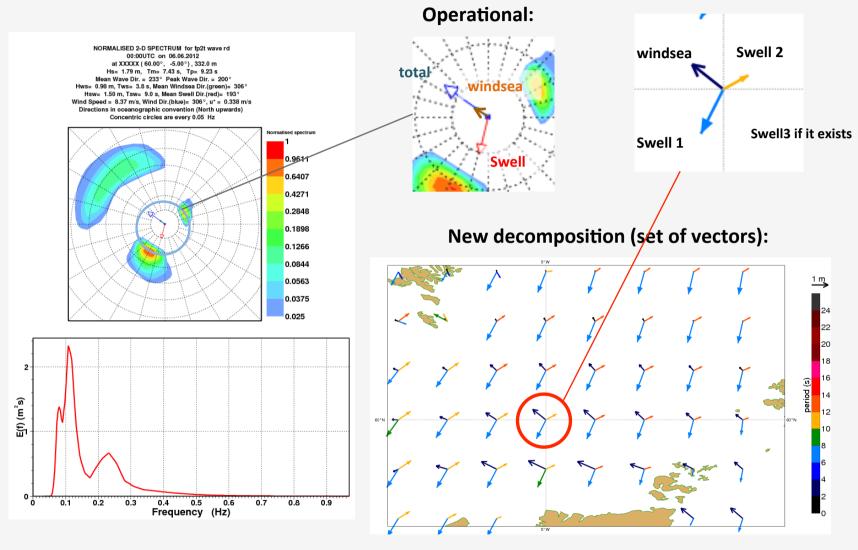


Windsea and swell: cross sea





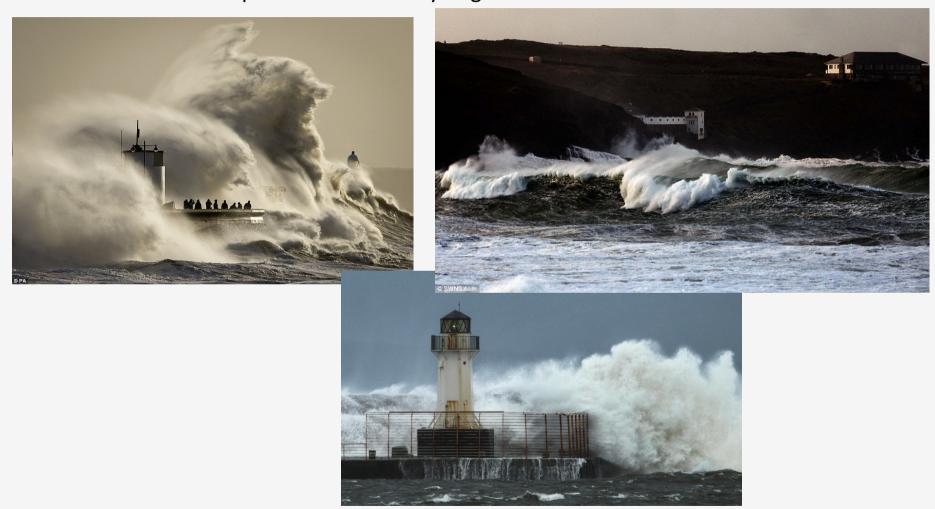
Since May 2015: spectral partitioning



arrow length: wave height,

arrow colour: corresponding mean wave period

At the end of December 2013 and beginning of January 2014, the UK and western Europe were battered by large waves:



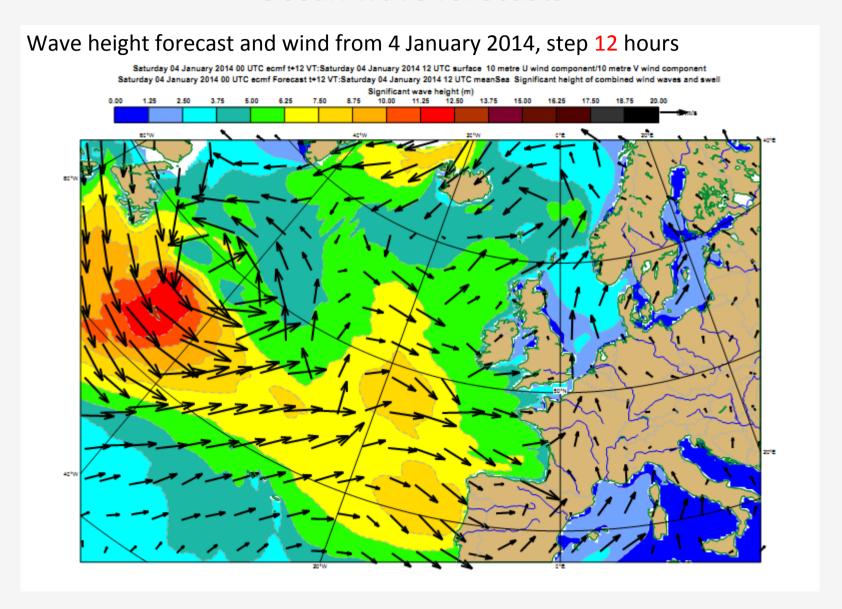
Then again in February and early March:

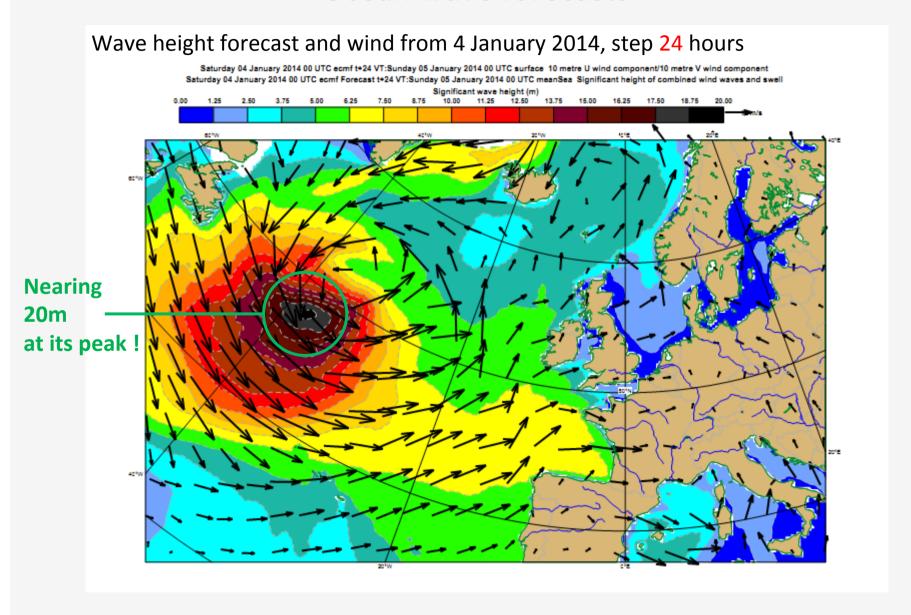


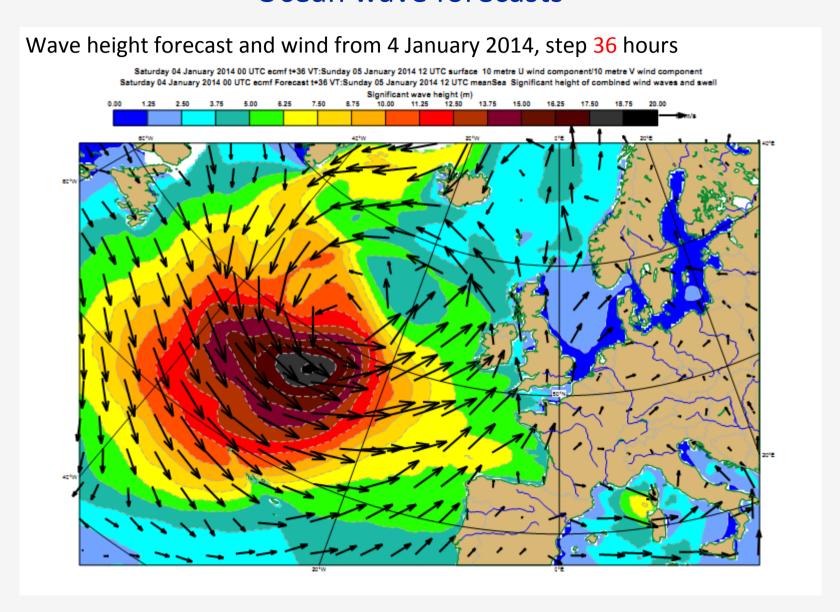
Porthleven Clock Tower, Cornwall

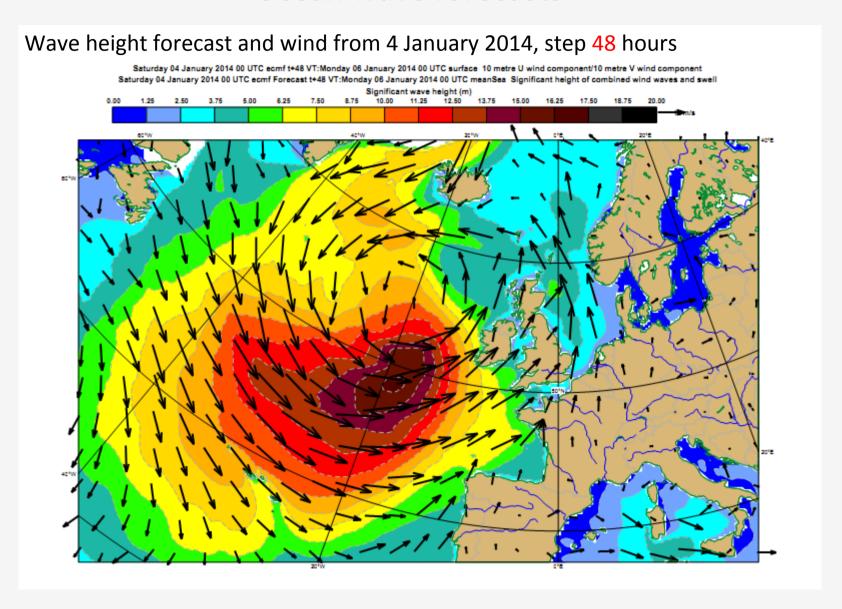
May 1, 2013

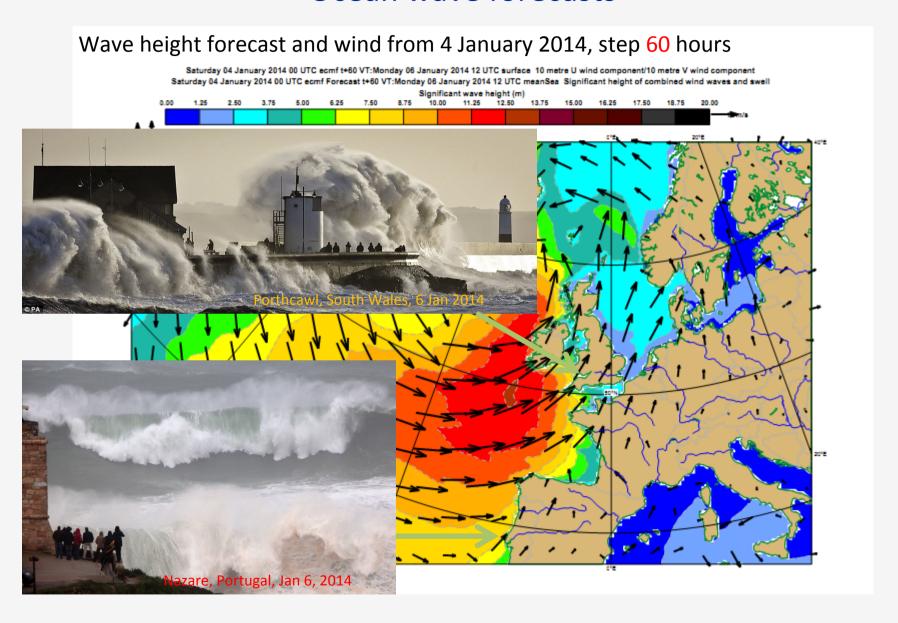
February 5, 2014









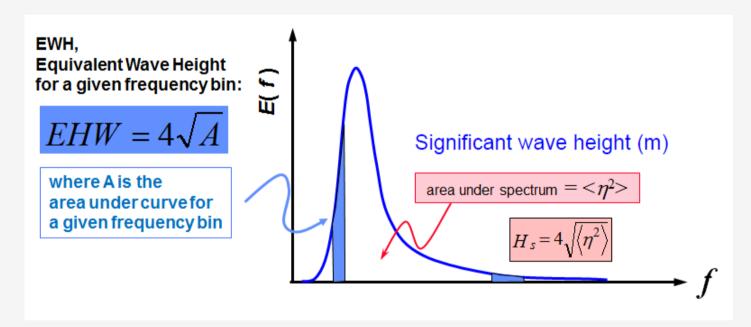


Long swell forecasts

Swell are long waves propagating away from storms.

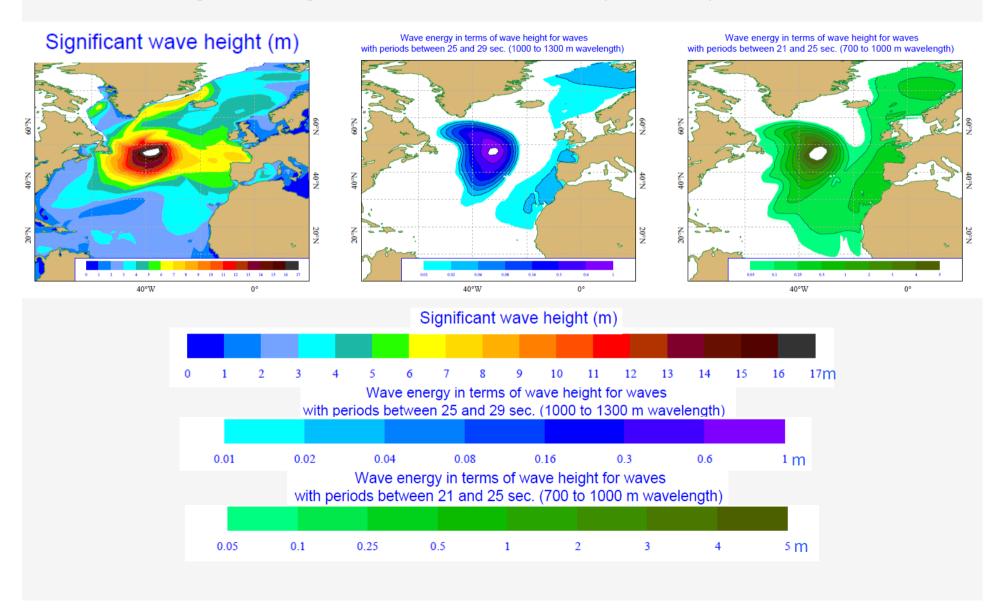
It is possible to follow the evolution of the swell.

Define the Equivalent Wave Height:



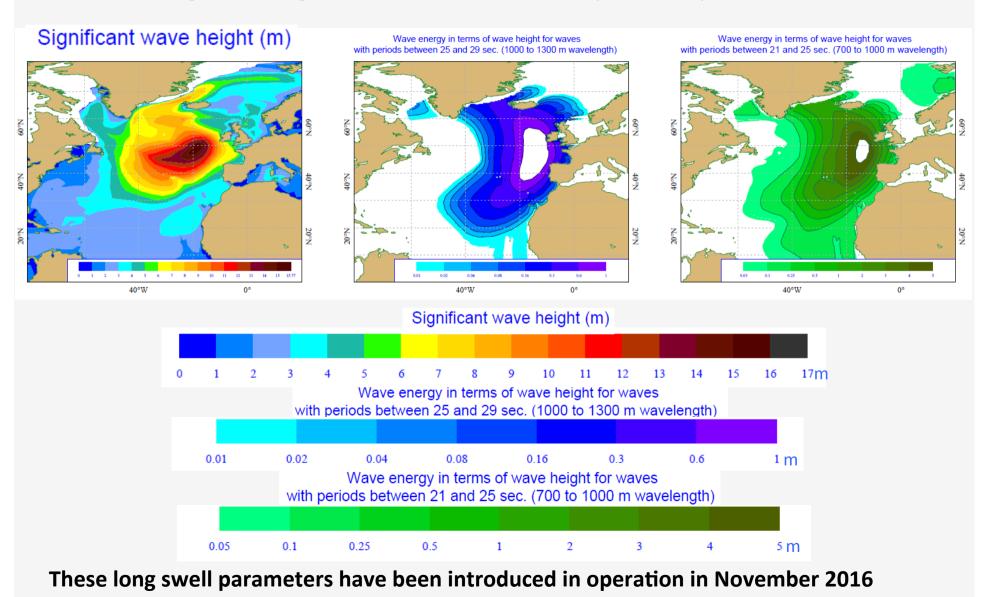
Long swell forecasts

Wave height and long swell forecast from 4 January 2014, step 24



Long swell forecasts

Wave height and long swell forecast from 4 January 2014, step 48

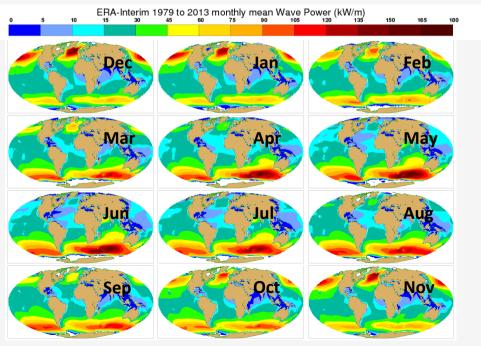


Wave energy:









Wave energy flux:

General definition

$$\Phi = \rho_{w} g \iint V_{g} F(f,\theta) df d\theta$$

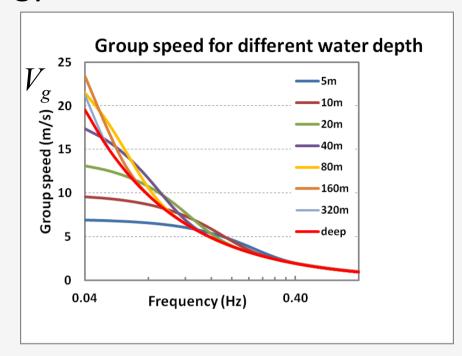
Its calculation would require integrating the spectra. But in deep water:

$$V_g = \frac{g}{4\pi f}$$

Define the spectral moments:

$$m_i = \iint f^i F(f,\theta) df d\theta$$

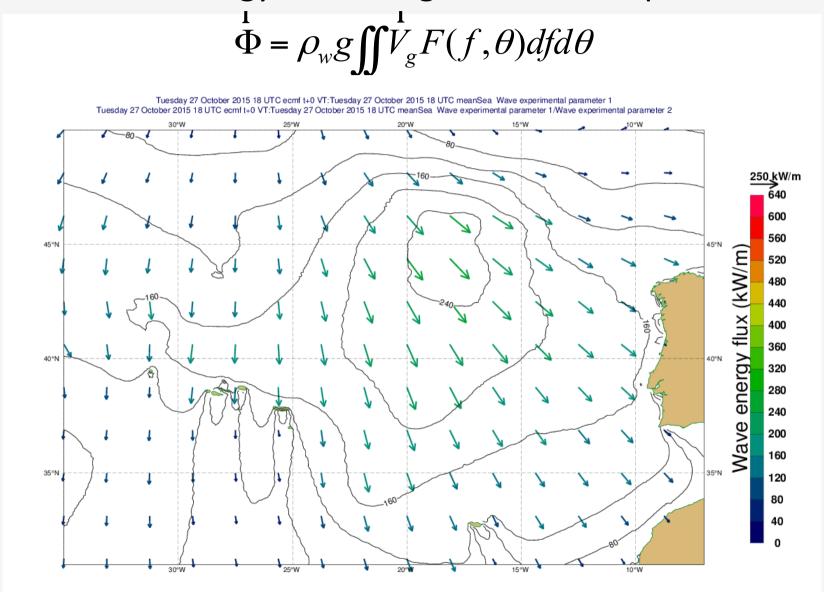
$$H_s = 4\sqrt{m_0}$$
 $T_m = \frac{m_{-1}}{m_0}$ Note that mean period T_m is based on the moment -1!



$$\Phi = \frac{\rho_w g^2}{64\pi} T_m H_s^2 \qquad \Phi \approx 0.478 T_m H_s^2 \text{ kW/m}$$
He and To are standard output of

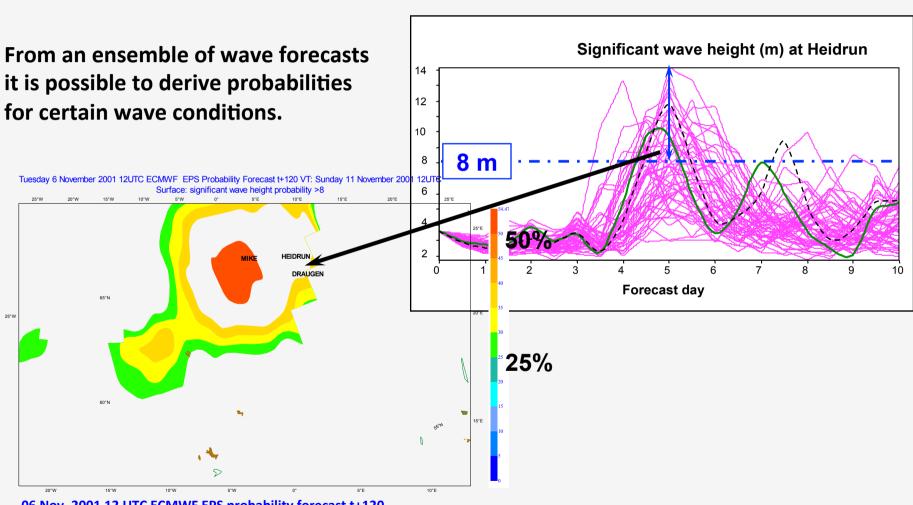
H_s and T_m are standard output of ECWAM

Wave energy flux: using archived 2d spectra



The magnitude and direction of the wave energy flux have been introduced in operation in November 2016

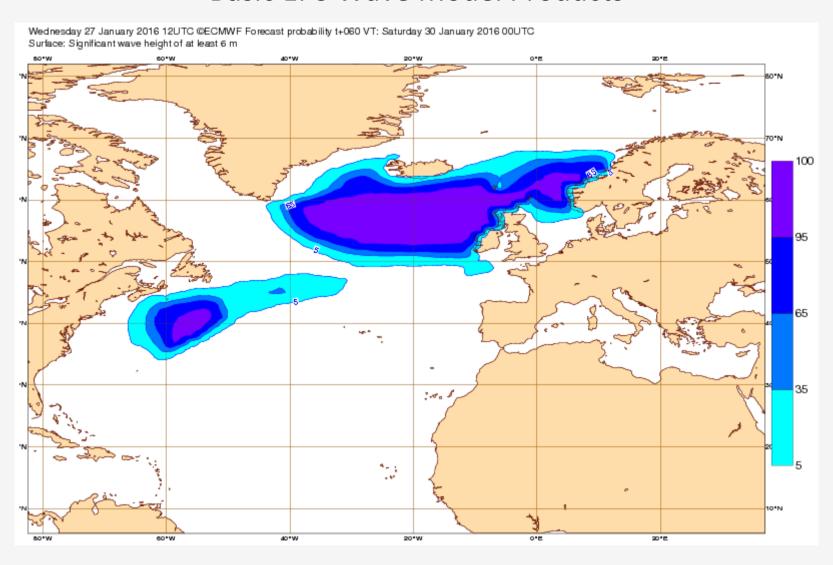
A more 'classic' use of the EPS:



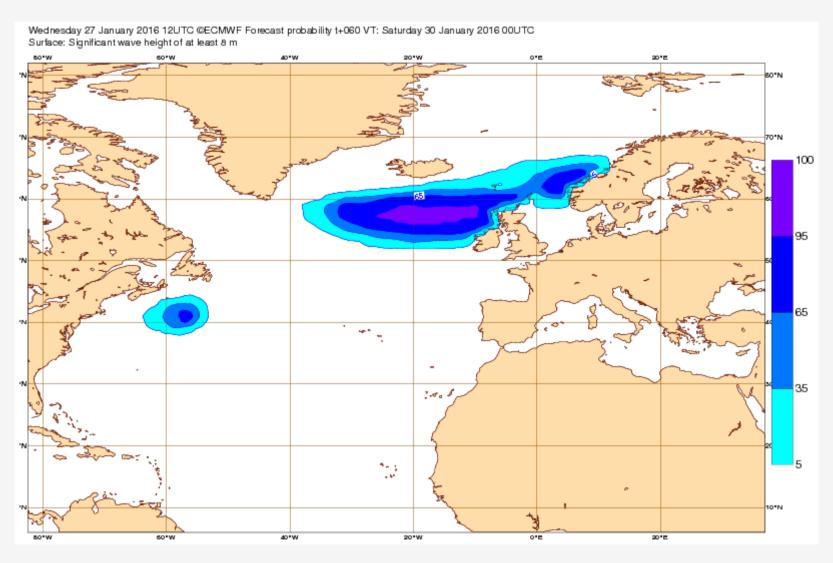
06 Nov. 2001 12 UTC ECMWF EPS probability forecast t+120

Significant wave height above 8 m

Basic EPS Wave Model Products



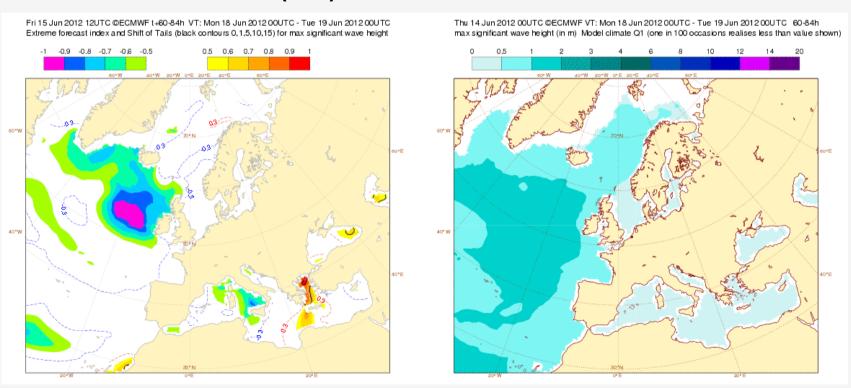
Basic EPS Wave Model Products



EFI plots

From the new model climate, it is possible to derive indices that indicate deviations in probabilistic terms from what is 'expected'.

Extreme Forecast Index (EFI): -1 means that all EPS are below climate.



EFI for significant wave height

01 percentile of the distribution for significant wave height

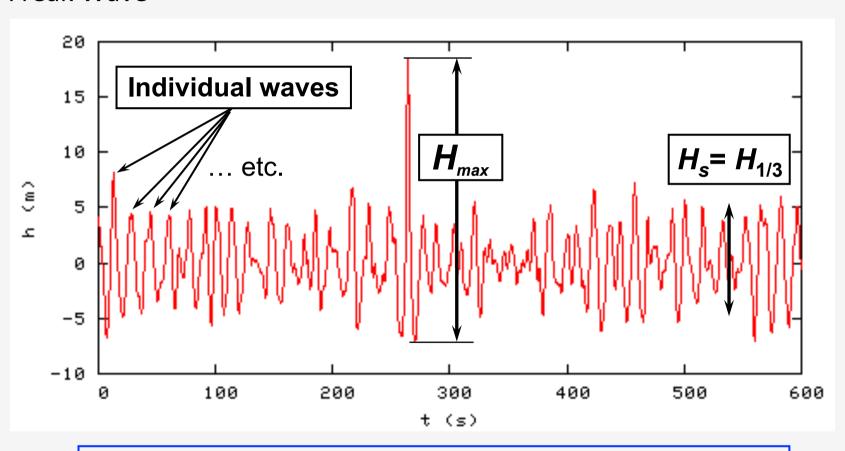
We are not always dealing with nice 'predictable' waves:







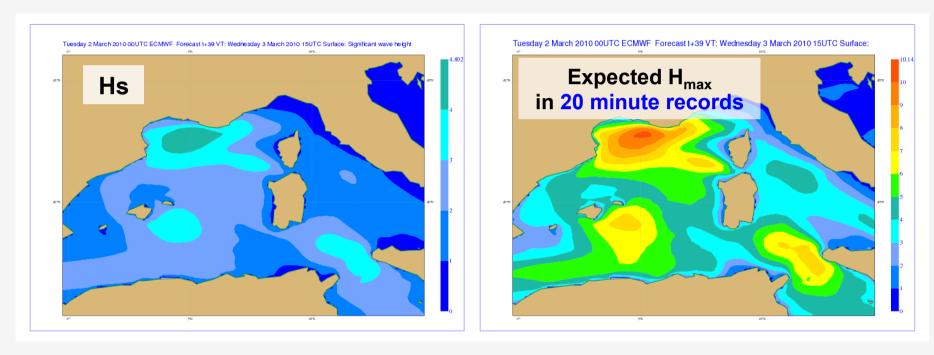
Individual Waves, Significant Wave Height, H_s , Maximum Individual Wave Height, H_{max} , and Freak Wave



If $H_{max} > 2.2 H_s$ freak wave event

Wave Model Products: Extreme Waves

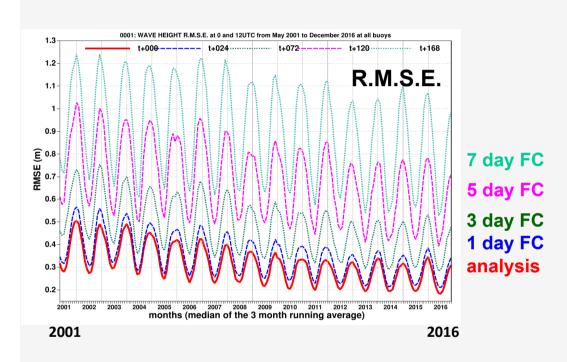
We have a parameter to estimate the height of the highest <u>individual</u> wave (H_{max}) one can expect. Its value can be derived from the 2d wave spectrum:



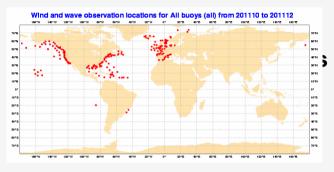
March 3, 2010, 15UTC Forecasts fields from Friday 2 March, 2010, 0 UTC

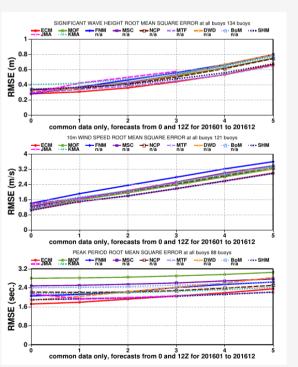
Continued general improvement of model forecasts For example: ECMWF forecast wave height against buoy measurements:

Home -> Forecasts -> Charts -> Verifications -> Wave Products Comparison



See also the Wave Forecast Verification Project maintained on behalf of the Expert Team on Waves and Storm Surges of the WMO-IOC Joint Technical Commission for Oceanography and Marine Meteorology (JCOMM)





http://www.jcomm.info/index.php?option=com content&task=view&id=131&Itemid=37

