Using OpenIFS to describe specific weather events at the Carpathian Basin

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OUTLINE

- 1. Motivation
- 2. Experimental design
- 3. Results
- 4. Conclusions, outlook

Background

- Hungarian Meteorological Service:
 - ALADIN/HARMONIE consortium, participation in ALADIN and AROME development
 - Operational weather forecasts based on ALADIN and AROME
 - Using ECMWF data and products (e.g., LBCs, EPS)
 - Teaching numerical modelling for meteorologist and mathematician students at the University (CHAPEAU, Lorenz model)
 - Coupled models (CHIMERE, etc.), regional climate modelling (ALADIN-Climate, REMO)
- Eötvös Loránd University:
 - WRF for everyday weather forecasts (GFS LBCs)
 - Adaptation of CHAPEAU model
 - Several coupled models (TREX, etc.)
 - Regional climate modelling based on 2 models

Motivation

- Two licence agreements for the OpenIFS model in Hungary:
 - Hungarian Meteorological Service (OMSZ; 2013)
 - Eötvös Loránd University (ELTE; 2014)
- Installation on the SGI supercomputer of OMSZ in 2013
- Cooperation of the two institutes on a Master Thesis: supervision of its synoptic part by University, its modelling part (including model execution) by OMSZ

Experimental design

- Resolution: T639 spectral truncation (~34 km horizontal resolution) and 91 vertical levels – full capacity of SGI
- Period: 13 November and 4 December 2011 with 24-hour integrations
- Spatial focus on Carpathian Basin
- Weather situation: persistent cold air pool

Cold air pool

- Persistent inversion in winter anticyclonic situation
- More humidity in the cold air layer → cooling down to its dew point temperature → fog during the night
- Fog does not dissolve in daytime → elevation → stratus layer
 → in daytime the cloud top radiates out → cooling
- Low mixing → dramatic impact on air pollution
- Frequent weather phenomenon over the Carpathian Basin in winter providing modelling challenge → great IFS development based on forecasters' feedbacks

Verification

- Location: Szeged
- Verification against surface and radiosonde measurements



- Evaluated variables:
 - mean sea level pressure
 - shallow convective potential energy (Bozóki, 1987)
 - top and base of inversion layer
 - height of cloud top and cloud base

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- Energy needed to shallow convection
- Area between state curve and dry adiabat from 850 hPa to surface
- Modification of CAPE equation (Williams & Renno, 1993):

$$SCP^{850} = -R_d \int_{p_c}^{850} (T_{vp} - T_{ve}) d(\ln p)$$

 T_{vp} and T_{ve} : virtual temperature for air parcel and environment

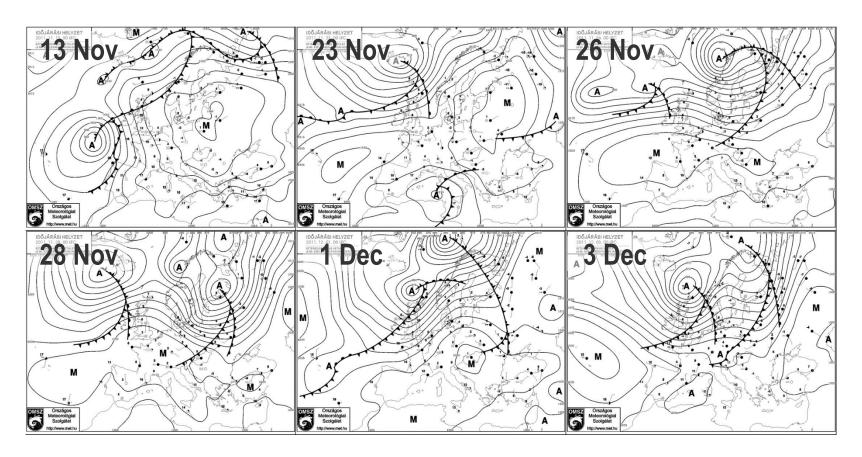
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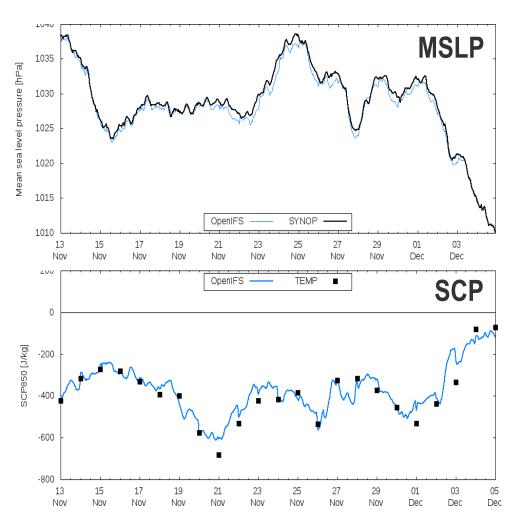
- Calculated based on vertical temperature gradient
- For measurements: from the surface
- For model data: using levels between 1000 and 700 hPa

Synoptic situation

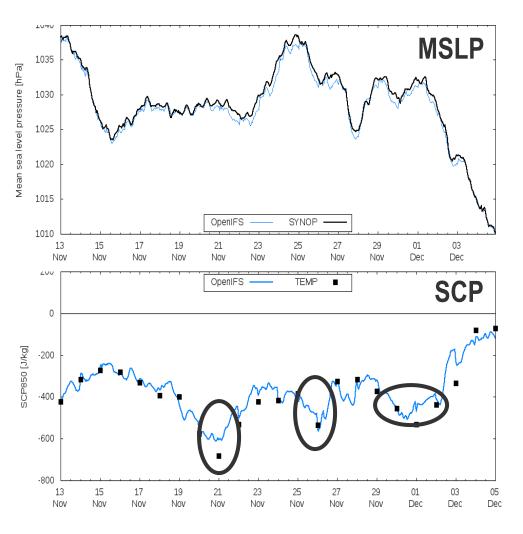


Anticylone over Eastern Europe (M) \rightarrow warm front \rightarrow broken cold air pool \rightarrow re-formation after cold front \rightarrow total break up

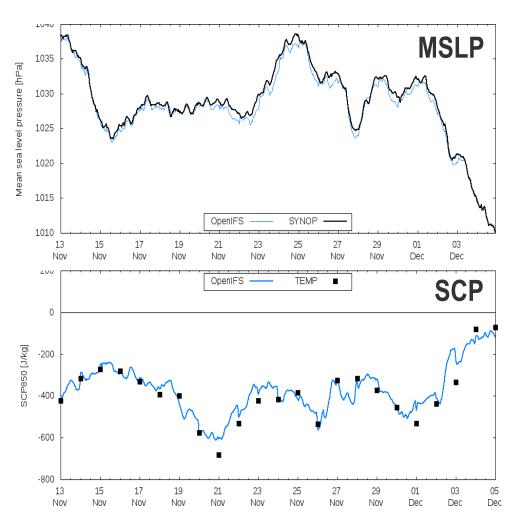
Results



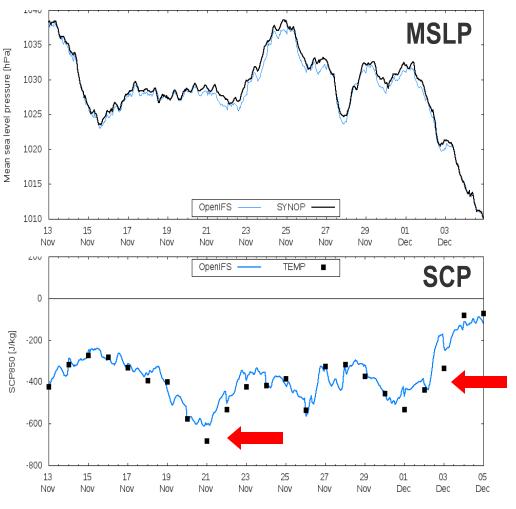
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- Negative SCP → stability in lower atmosphere
- 3 SCP minima
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 → slower development of cold air pool



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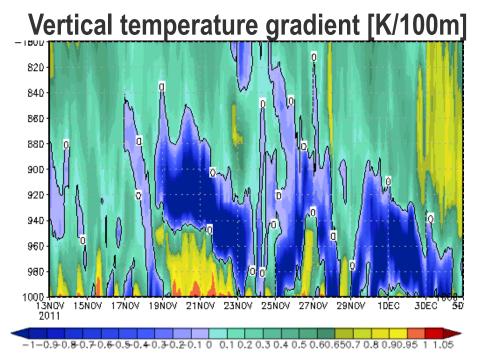


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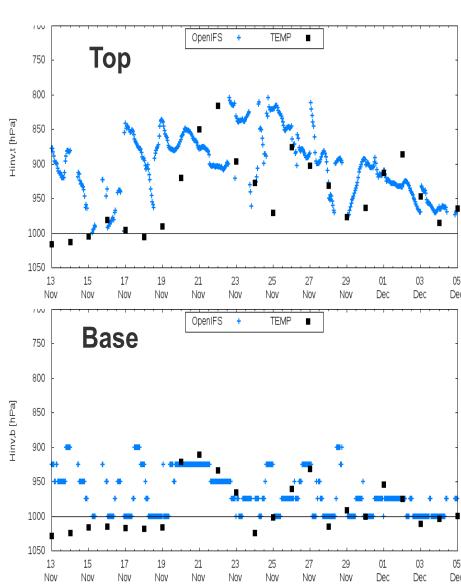


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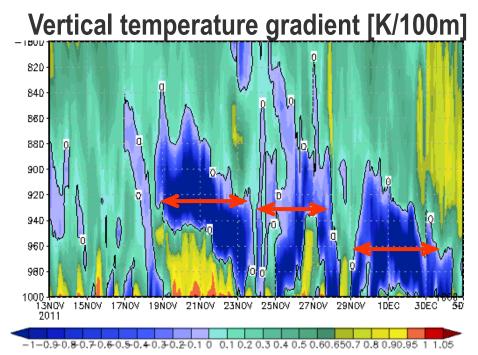
Inversion layer



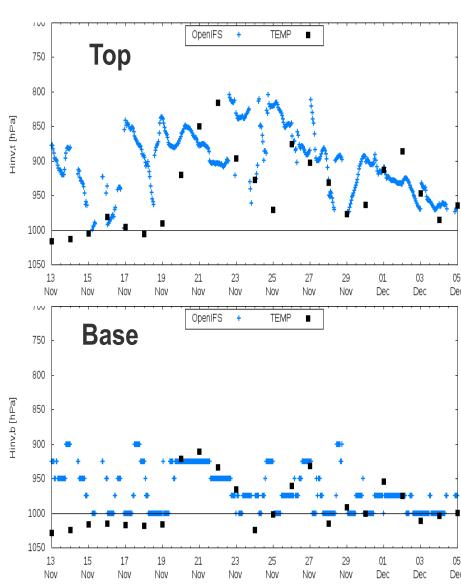
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- Base: good during cold air pool



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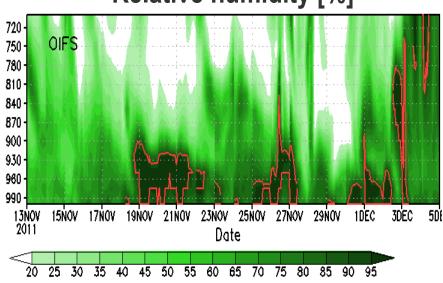


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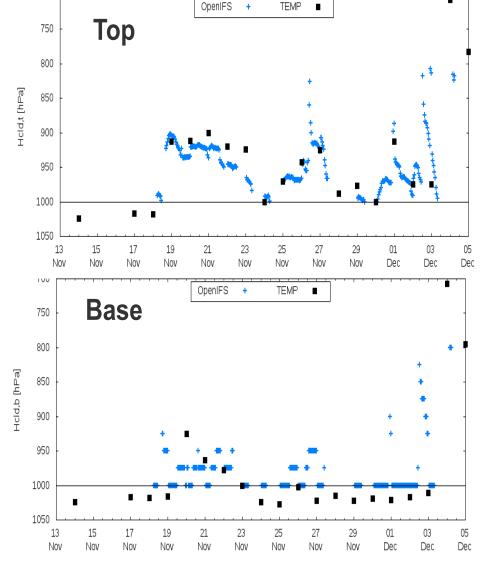


Cloud top and base





- Good RHU results during cold air pool events
- Fog was not well reproduced (due to some technical issues)

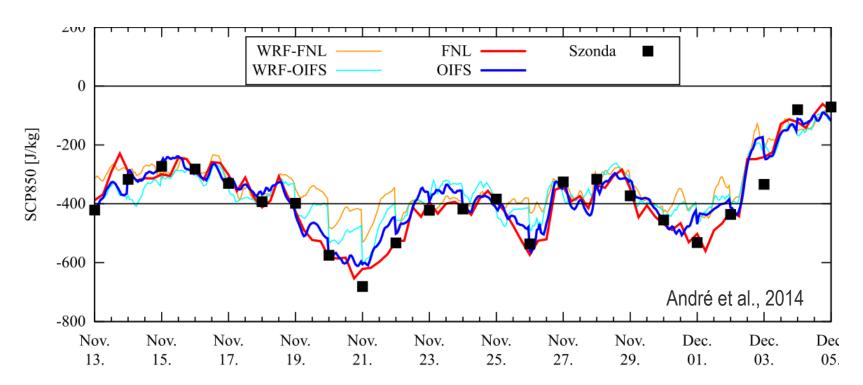


Conclusions

- Using SCP⁸⁵⁰ an appropriate definition can be created by the persistent cold air pool → search algorithm
- Relative humidity was not a good precursor of the phenomenon
- OpenIFS was capable of describing persistent cold air pool

Outlook

- OpenIFS provided LBCs for limited area WRF runs
- Comparison of WRF runs driven by GFS analyses



Outlook

Thank you for your attention!

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